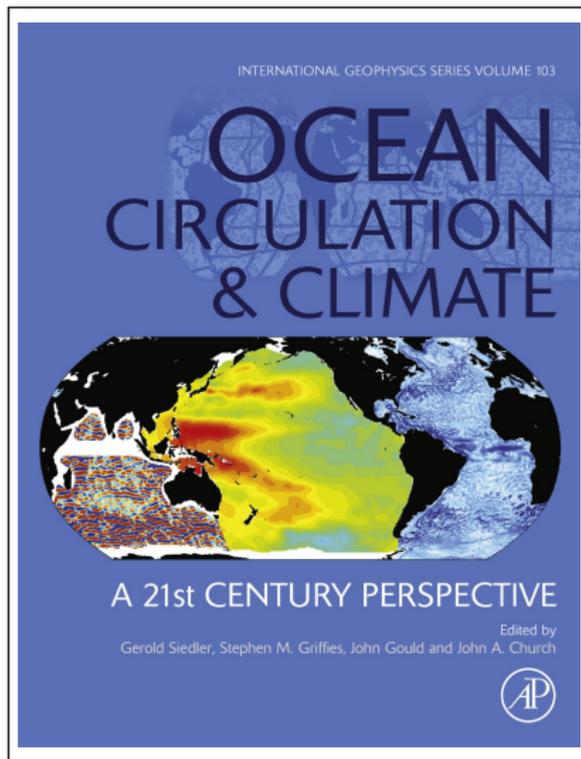
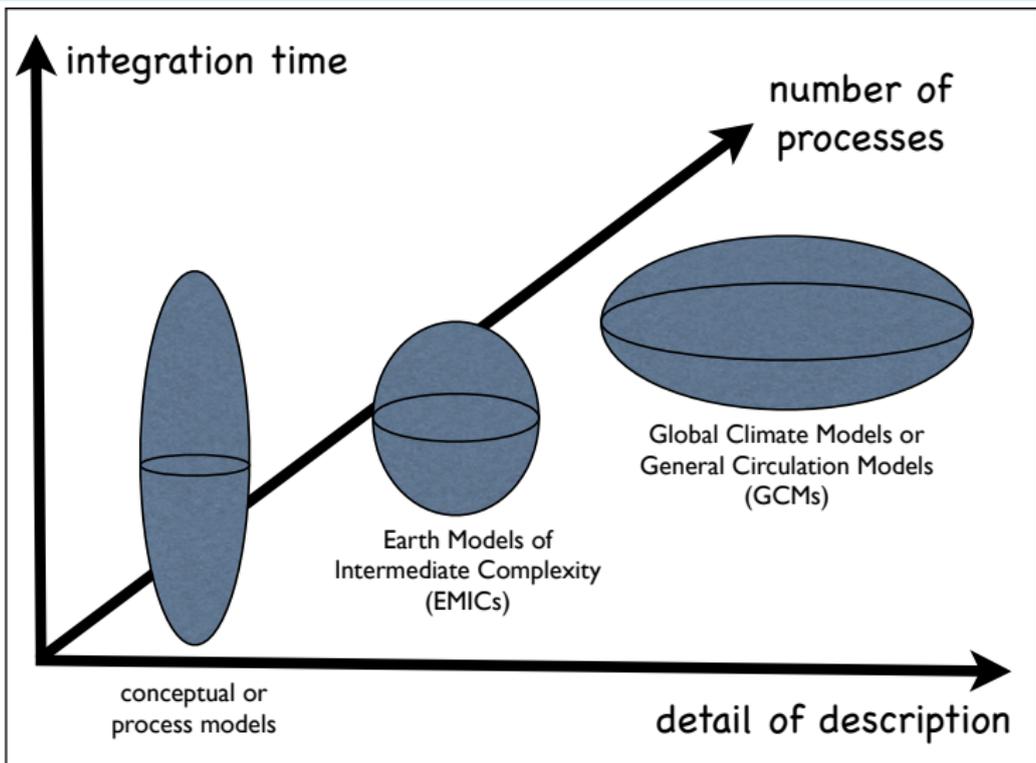


Ocean Circulation and Climate, 2nd Edition (2013)



Types of ocean and climate models

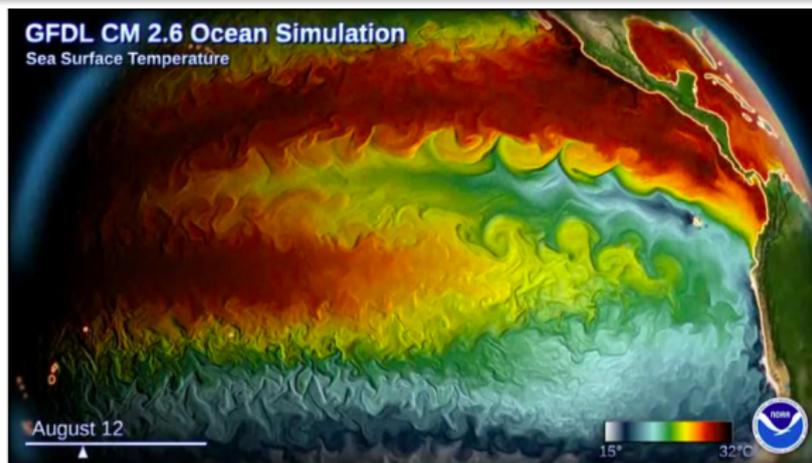


Compliments of Stephanie Waterman, University of British Columbia, Canada

SST animation from GFDL CM2.6 climate model

Animation 1: Daily SST from the GFDL CM2.6

This coupled climate model uses a 0.1° configuration of MOM5 for the ocean component, under a 50 km global atmosphere model. It has been integrated for multiple-centuries in support of climate and ocean related studies. Available from [Vimeo](#)





Equations for ocean dynamics and thermodynamics

Thermo-hydrodynamic equations for the ocean

Momentum (Newton's 2nd Law)

$$\rho [\partial_t + (2\Omega + \omega) \times] v = -\rho \nabla (KE + GPE) + \nabla \cdot (\tau - 1p)$$

Mass conservation (continuity)

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho v) = 0$$

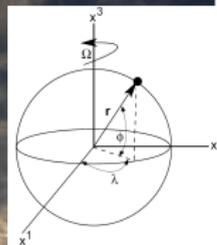
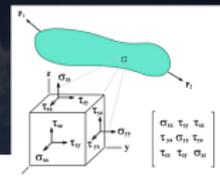
Enthalpy (heat) conservation

$$\frac{\partial(\rho\theta)}{\partial t} + \nabla \cdot (\rho v \theta + J_\theta) = 0$$

Salt conservation

$$\frac{\partial(\rho S)}{\partial t} + \nabla \cdot (\rho v S + J_S) = 0$$

Equation of state relates density to temp, salinity, pressure $\rho = \rho(\theta, S, p)$



Key dynamical speeds

Acoustic: $C_s \sim 1500 \text{ m s}^{-1}$

External gravity: $\sqrt{gH} \sim 150 \text{ m s}^{-1}$

Internal gravity: $NH \sim 3 \text{ m s}^{-1}$

Advection: $U \sim 1 \text{ m s}^{-1} \ll \#$

7/25/16

Griffies talk to IITM in Pune, India





A sample of ocean processes

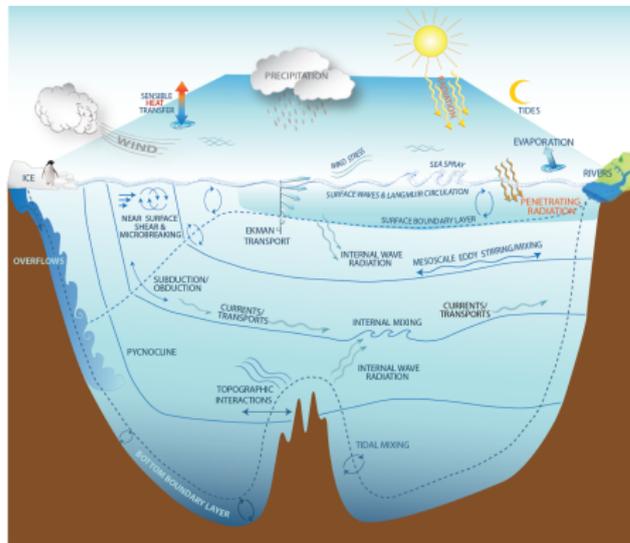
Slides in this section

- A zoo of physical ocean processes
- Space-time diagram of ocean motions
- Upper ocean boundary and wave interactions
- The marginal ice zone (MIZ)
- Southern Ocean processes
- The value of idealized Southern Ocean simulations
- Turbulent cascade of mechanical energy





A zoo of physical ocean processes

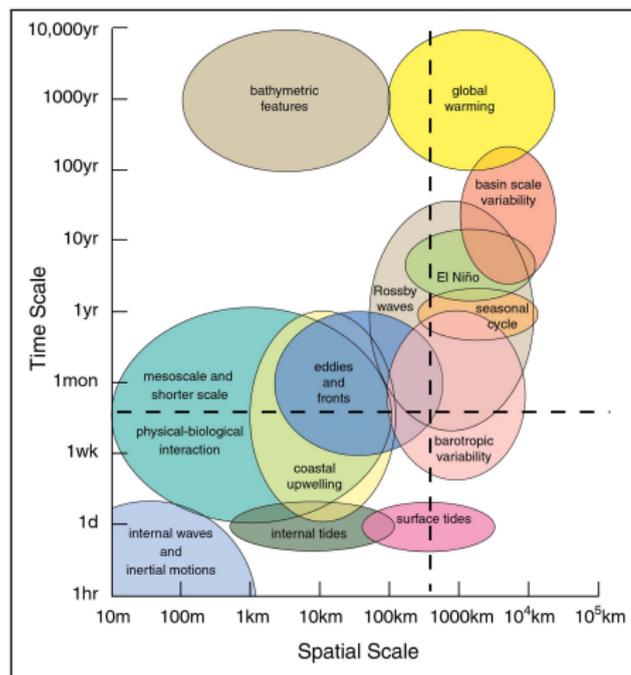


- The ocean contains a zoo of physical processes!
- Strong coupling between processes \Leftrightarrow no spectral gap.
- Coupling means it is generally better to resolve than parameterize.
- Yet we cannot resolve everything \Rightarrow a practical need for parameterizations that pass the “laugh test”.

From Griffies and Treguer (2013)



Space-time diagram of ocean dynamical processes



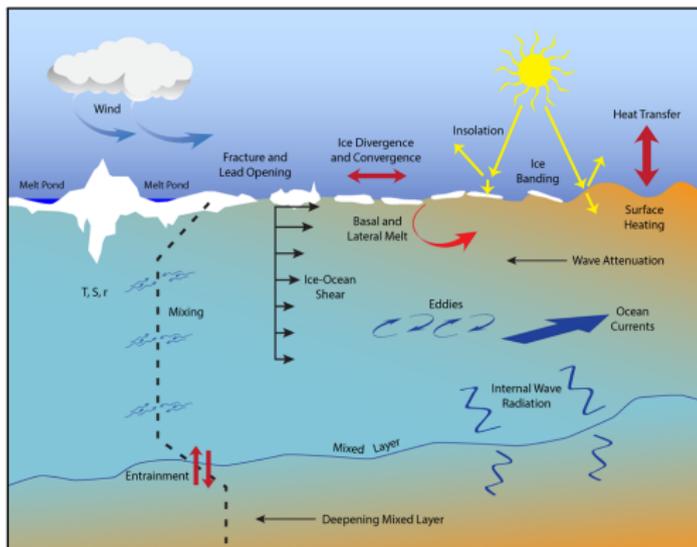
- Broad range of space-time scales
- We again see the absence of a clear spectral gap except for scales larger than 1000 km.

From Chelton (2001)





The marginal ice zone (MIZ)

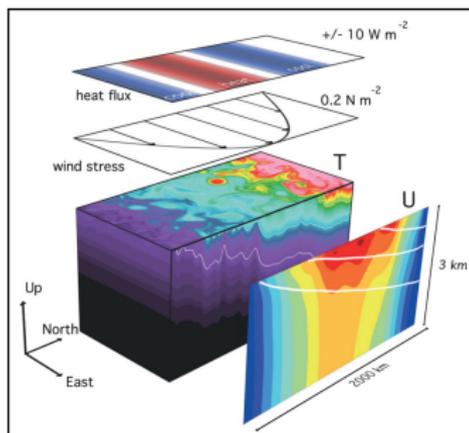


From [ONR Marginal Ice Zonal Project](#)

Questions about processes at the marginal ice zone are of prime importance as Arctic sea ice melts.



A Southern Ocean process study



From Abernathy et al. (2011)

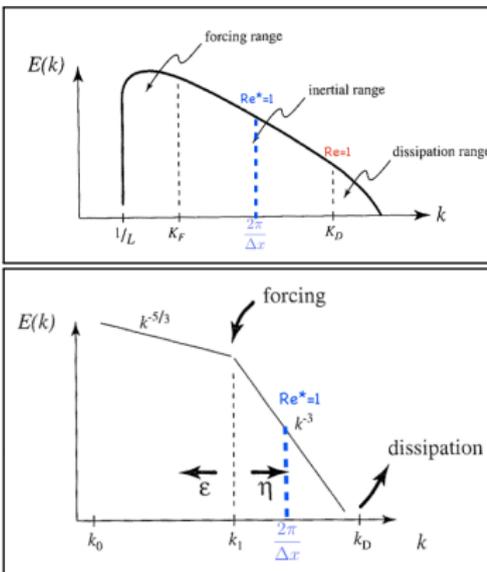
- The Southern Ocean is a region where mesoscale eddies are of leading order importance.
- This animation is part of an idealized study, and is shown here as an example how idealized process models can lend useful insight into the real ocean.

Animation 3: Southern Ocean channel

Animation from R. Abernathy, available from [Vimeo](#)



Turbulent cascade of mechanical energy



- 3d turbulence: energy cascade to small scales
- 2d/QG turbulence: energy cascade to large scales (inverse cascade)
- Cascades act to couple space-time scales.

Compliments of Baylor Fox-Kemper, Brown University, USA

Animation 4: QG turbulence cascade

Compliments of Shafer Smith, NYU USA



The ocean parameterization problem

Slides in this section

- Resolving versus parameterizing: some numbers
- Facets of what we mean by “resolution”
- Spatial scale of mesoscale and submesoscale eddies
- Resolution required to represent mesoscale eddies
- Ocean resolution in IPCC-class climate models



Resolving versus parameterizing: some numbers

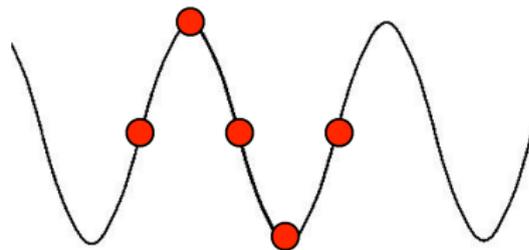
- Direct Numerical Simulation (DNS) of global ocean climate requires 3×10^{10} time steps of one second (1000 years).
- Setting the model's grid scale to the Kolmogorov length $\Delta = 10^{-3}\text{m}$ over a global ocean domain of volume $1.3 \times 10^{18}\text{ m}^3$ requires 1.3×10^{27} discrete grid cells. This is roughly $10^4 \times$ Avogadro's number!
- Each model grid point has a velocity vector and tracer fields to time integrate.
- Conclude:
 - We will be dust long before DNS of global ocean climate.
 - We must use parameterizations to simulate the ocean.
 - The rationalization of a DNS simulation typically requires a coarse-grained perspective, as certainly would DNS of the World Ocean climate.





Facets of what we mean by “resolution”

- general principles of resolution are the same for both atmospheric and ocean models
- there are different rules of thumb: one is that it takes 5 grid points to accurately define a feature without aliasing
- this means **1/8°** global resolution with an average horizontal grid cell of **14 km** can accurately depict only features larger than **56 km**
- models with variable grid spacing have variable resolution - beware of resolution-dependent physics!
- resolution is not cheap - because of the CFL* condition, as we shrink the horizontal grid spacing we must add vertical layers and decrease the time step



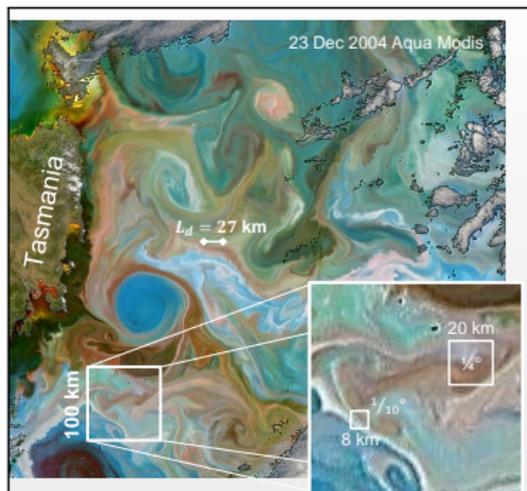
“every halving of the grid spacing requires roughly ten times as many computations”

* no transport faster than one grid cell per time step!

Compliments Stephanie Waterman, University of British Columbia, Canada



Spatial scale of mesoscale and submesoscale eddies



MODIS satellite w/ inserts by A. Adcroft (GFDL)

- Eddy size \propto first baroclinic Rossby Radius $\lambda_m = c_m/|f|$, where the phase speed is approximated by (Chelton et al. 1998)

$$c_m \approx \frac{1}{m \pi} \int_{-H}^0 N dz.$$

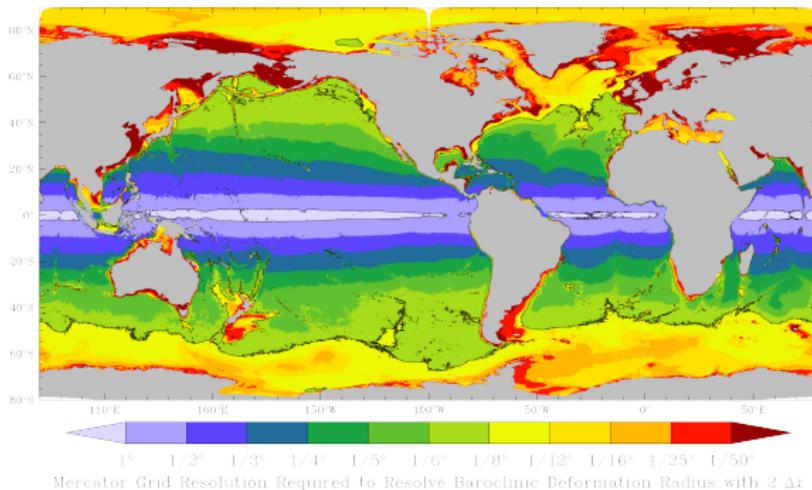
- Global models are marginal at representing this scale; regional and process models just reach into the submesoscale.

Animation 5: Southern Ocean regional process model

MITgcm w/ $1/20^\circ$ (and $1/80^\circ$ local refinement) w/ 150 vertical levels.
Available from [YouTube](#)



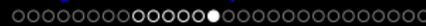
Resolution required to represent mesoscale eddies



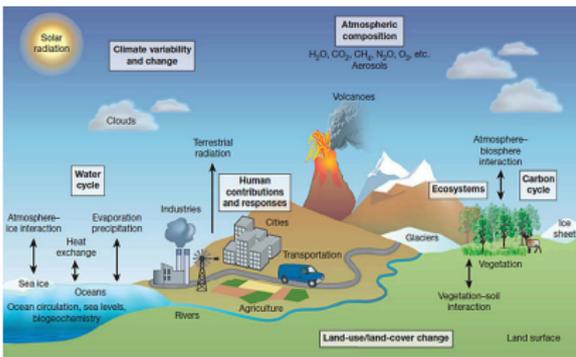
From Hallberg (2013)

- Hallberg (2013): $2\Delta \leq \lambda_1$ needed to resolve mesoscale eddies.
- Map indicates the necessary Mercator spacing for $2\Delta = \lambda_1$.

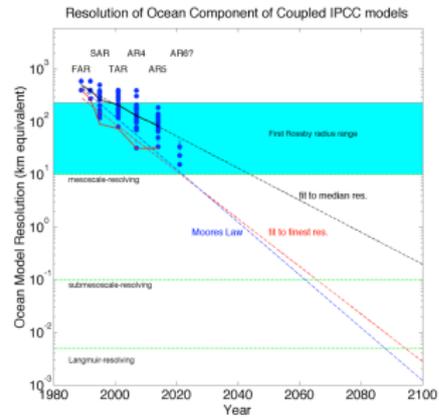




Ocean resolution in IPCC-class climate models



Compliments of GFDL



From B. Fox-Kemper, Brown University, USA

- The ocean is but one component amongst many within climate system models.
- Resolution refinement is painfully slow!
- This diagram is useful to target one's career choices.



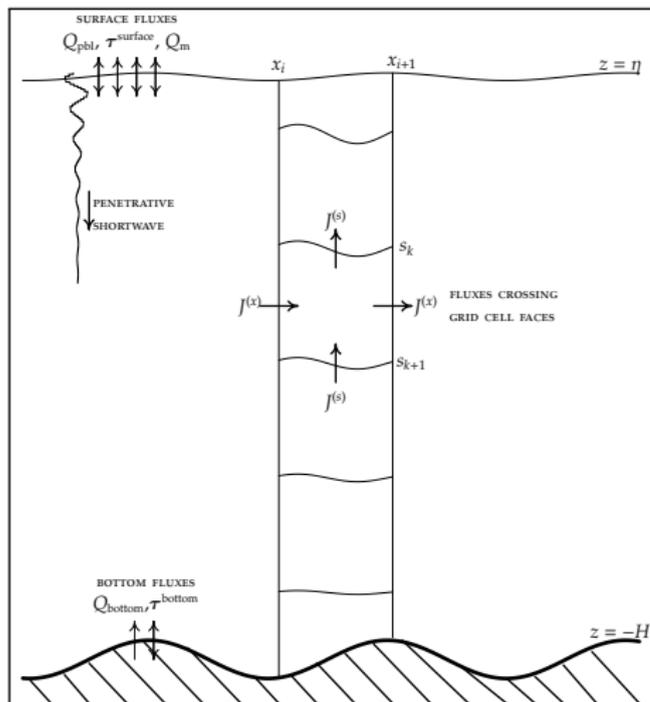
Partitioning the vertical

Slides in this section

- Discretizing a column of ocean fluid
- Vertical coordinate representation
- Geopotential or pressure vertical coordinates
- Isopycnal vertical coordinates
- Sigma or terrain following vertical coordinates



Discretizing a column of ocean fluid



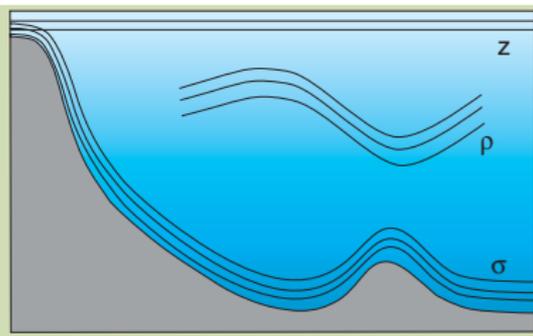
- Boundary fluxes through surface and bottom.
- Transport convergence (advective and subgrid scale), body forces (gravity, Coriolis), contact forces (pressure, friction), and penetrative radiation render time tendency for mass, tracer, and momentum.
- Generally fix the horizontal position of grid cells, but allow for upper and lower interfaces to be functions of time (e.g., z^* , pressure, σ , isopycnal)

From Griffies and Treguer (2013)



Vertical coordinate representation

Schematic of an ocean basin illustrating the three regimes of the ocean germane to the considerations of an appropriate vertical coordinate. The surface mixed layer is naturally represented using fixed-depth z (or pressure p) coordinates, the interior is naturally represented using isopycnal ρ (density tracking) coordinates; and the bottom boundary is naturally represented using terrain-following σ coordinates (after Griffies et al., 2000).

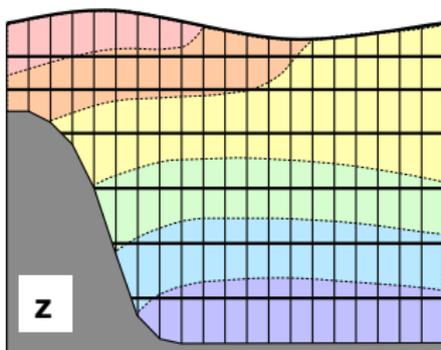


Adapted by Chassignet et al (2006) from original figure in Griffies et al (2000)

- GEOPOTENTIAL OR PRESSURE: common for non-hydrostatic process modelling and large-scale climate modelling ([MITgcm](#), [MOM](#), [NEMO](#))
- ISOPYCNAL: clean representation of interior quasi-adiabatic flows and overflows ([GOLD](#), [HYCOM](#))
- SIGMA OR TERRAIN FOLLOWING: common for shelf/coastal modelling (([ROMS](#)) and Curchitser's lectures)



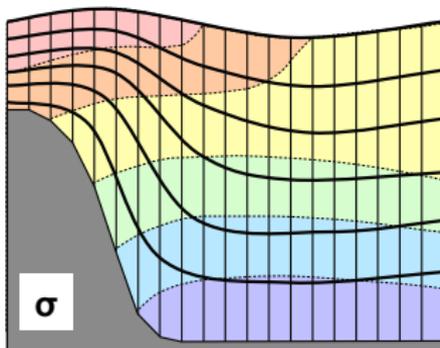
Geopotential or pressure vertical coordinates



- Most common method for global models; extensive experience.
- Generalizations: $z^* = H(z - \eta)/(H + \eta)$ absorbs SSH undulations; pressure (mass conserving).
- Spurious diapycnal mixing if poorly chosen numerical methods & parameter settings (e.g., Ilicak et al 2012).
- Downslope flows poorly represented absent very fine resolution (Winton et al. 1998).



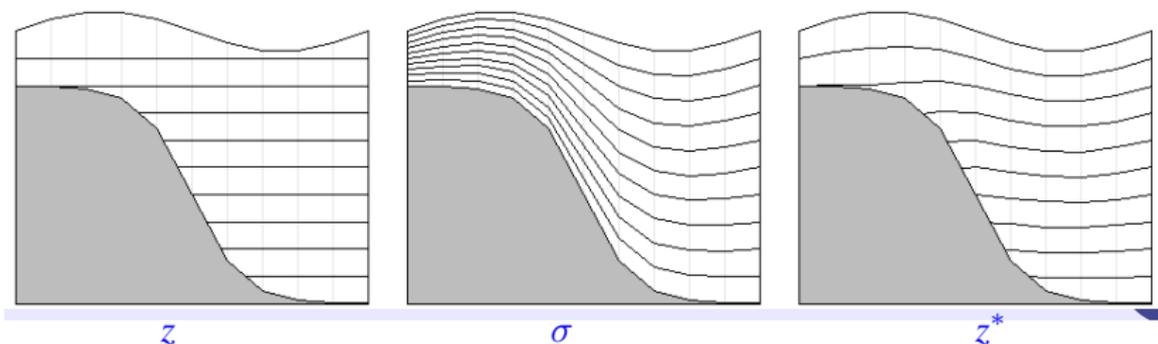
Sigma or terrain following vertical coordinates



- Extensive applications for coasts & shelves
- Traditionally $\sigma = (z - \eta)/(H + \eta)$, but with generalizations.
- As for geopotential, \exists spurious diapycnal mixing with poorly chosen numerical methods & parameter settings.
- Much care is needed to handle horizontal pressure gradient calculation; generally need to smooth topography.
- There are very few global climate realizations.



Undulations of coordinate surfaces w/ z , σ , and z^*

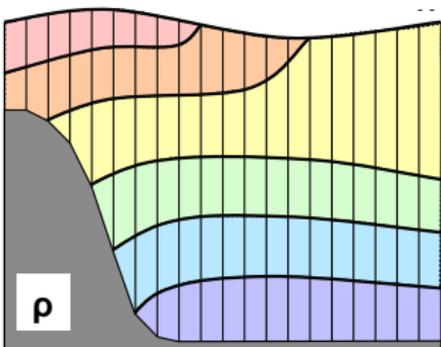


Animation 6: undulations of coordinate surfaces

Animation to illustrate undulating z , σ , and z^* coordinate surfaces in the presence of a gravity wave. Compliments A. Adcroft (GFDL).



Isopycnal vertical coordinates



- Quasi-adiabatic interior & flow-topography interactions (e.g., overflows)
- Inherently poor representation if weak vertical stratification (e.g., Labrador Sea, Southern Ocean, coastal regions).
- Care needed to represent realistic ocean thermodynamics and conservative transport, though proven methods now common.
- GFDL-GOLD, HYCOM, and Bergen



Partitioning the horizontal

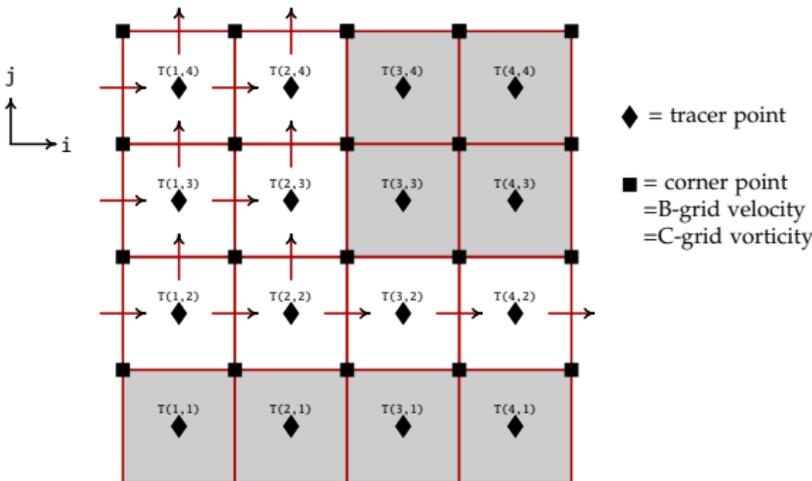
Slides in this section

- Horizontal representation: structured finite volume
- Examples of structured finite volume grids
- Horizontal representation: unstructured finite volume
- Example of unstructured finite volume grid
- Horizontal representation: unstructured finite element
- Examples of unstructured finite element meshes



Horizontal representation: structured finite volume

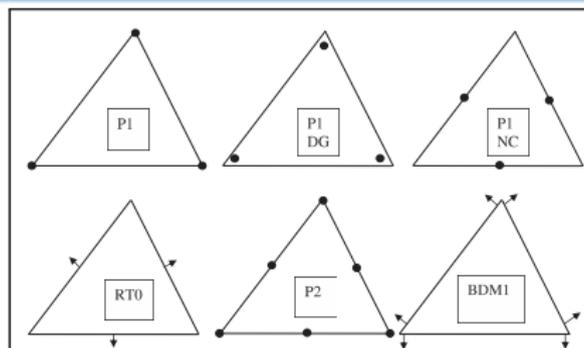
TRACER CELLS $T(i, j)$ WITH FLUXES AND LAND/SEA MASKING



- Most common approach since 1960s; e.g., [HYCOM](#), [MITgcm](#), [MOM](#), [NEMO](#), [ROMS](#).
- Recent advances with nesting allow for multi-scale simulations (Debreu and Blayo, 2008)



Horizontal representation: unstructured finite element



From Danilov (2013)

- Aimed at seamless representation of multiple-scales.
- Decompose continuous equations using basis functions and matrix inversions.
- Indirect addressing adds computational cost.
- Effort at AWI focused on climate: [FESOM](#)
- Effort at Louvain focused on shallow ocean: [SLIM](#)
- Non-hydrostatic process model at Stanford: [SUNTANS](#)

